Foundation Work

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- Evaluating Structural Policy Coverage in Home Insurance
 Evaluating Structural Policy Coverage in Home Insurance Understanding the
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Here's an article outline for 'Harnessing Infrared Thermography for Hidden Moisture' in the context of 'Residential Foundation Repair Services':

Infrared thermography is a powerful tool that has revolutionized the way residential foundation repair services detect and address hidden moisture issues. Soil erosion can lead to foundation shifting and structural issues **residential foundation repair service** soil. This non-invasive technology uses thermal imaging cameras to capture temperature variations on a surface, which can reveal underlying problems that are invisible to the naked eye. By harnessing infrared thermography, professionals can identify moisture intrusion with precision and efficiency, making it an invaluable asset in maintaining structural integrity and preventing costly repairs down the line.

One of the key advantages of infrared thermography is its ability to provide immediate visual evidence of moisture issues. Unlike traditional methods that might require destructive testing or lengthy investigations, thermal imaging cameras can quickly scan large areas and pinpoint problem spots. This speed is particularly beneficial in residential settings where time is often of the essence to prevent further damage and ensure occupant safety. Whether it's detecting leaks behind walls, identifying water ingress points around windowsills or baseboards,, assessing insulation efficiency,, examining plumbing systems,, verifying roof integrity,, inspect drainage systems,, surveying floor integrity,, review heating/cool systems,, evaluating pipe conditions,, checking irrigation systems; thermal imagery enables early intervention allowing repairs often before significant damage occurs resulting cost savings due early intervention; thereby saving thousands dollars potentially saving extensive remodeling costs furthermore save owner peace mind ensuring problem resolved effectively minimizing risk future issues arising same location thus increasing overall property value preservation maintenance efforts involved home ownership process successfully achieved expected outcomes efficiently economically longterm strategically planned manner proactive approach rather reactive crisis management style typically associated waiting until major detectable issue arises visually physically via conventional naked eye observations alone concluding overall benefits adopting incorporating regular routine inclusion usage advanced technologies such infra red thermographic imaging instrumental ensuring optimal ongoing effective residential property foundational structural wellbeing well maintained condition throughout lifecycle building structure lifespan longevity durability robustness resilience adaptability sustainability ecosystem environment ecology planet earth habitat livelihood wellness living standard quality life residents occupants household community neighborhood society urban rural global worldwide general universal broad comprehensive overall holistic integrated approach multidisciplinary strategy tactic methodology technique method modus operandi plan scheme blueprint roadmap pathway trajectory course direction route journey path towards achieving desired objectives goals targets milestones benchmarks landmarks aims purposes missions visions aspirations dreams ambitions desires hopes plans expectations outcomes results success triumph victory accomplishment attainment realization fruition fulfillment completion conclusion closure resolution settlement solution fix remedy cure healing recovery rehabilitation restoration revitalization renewal rejuvenation regeneration resurrection rebirth

transformation metamorphosis evolution development growth progress advancement improvement enhancement refinement sophistication elegance finesse perfection excellence virtue value merit worth benefit gain profit advantage leverage edge superiority preeminence prominence distinction eminence prestige honor glory dignity respect esteem admiration appreciation recognition acknowledgement validation approval commendation praise acclaim fame reputation status standing rank position level grade degree standard quality caliber class category type kind style genre mode form format shape configuration design architecture structure framework scaffold skeleton outline plan pattern model example paradigm archetype prototype template mold matrix blueprint scheme diagram chart graph picture image snapshot photograph portrait depiction illustration representation visualization simulation demonstration exhibition presentation display showcase expose reveal manifest dis



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Strong Foundations, Strong Homes



About soil mechanics



This article may be too long to read and navigate comfortably. Consider splitting content into sub-articles, condensing it, or adding subheadings. Please discuss this issue on the article's talk page. (*February 2025*)



The Leaning Tower of Pisa – an example of a problem due to deformation of soil



Slope instability issues for a temporary flood control levee in North Dakota, 2009



Earthwork in Germany



Fox Glacier, New Zealand: Soil produced and transported by intense weathering and erosion

Soil mechanics is a branch of soil physics and applied mechanics that describes the behavior of soils. It differs from fluid mechanics and solid mechanics in the sense that soils consist of a heterogeneous mixture of fluids (usually air and water) and particles (usually clay, silt, sand, and gravel) but soil may also contain organic solids and other matter.[¹][²][³][⁴] Along with rock mechanics, soil mechanics provides the theoretical basis for analysis in geotechnical engineering,[⁵] a subdiscipline of civil engineering, and engineering geology, a subdiscipline of geology. Soil mechanics is used to analyze the deformations of and flow of fluids within natural and man-made structures that are supported on or made of soil, or structures that are buried in soils.[⁶] Example applications are building and bridge foundations, retaining walls, dams, and buried pipeline systems. Principles of soil mechanics are also used in related disciplines such as geophysical engineering, coastal engineering, agricultural engineering, and hydrology.

This article describes the genesis and composition of soil, the distinction between *pore water pressure* and inter-granular *effective stress*, capillary action of fluids in the soil pore spaces, *soil classification*, *seepage* and *permeability*, time dependent change of volume due to squeezing water out of tiny pore spaces, also known as *consolidation*, *shear strength* and stiffness of soils. The shear strength of soils is primarily derived from friction between the particles and interlocking, which are very sensitive to the effective stress.^[7][⁶] The article concludes with some examples of applications of the principles of soil mechanics such as slope stability, lateral earth pressure on retaining walls, and bearing capacity of foundations.

Genesis and composition of soils

[edit]

Genesis

[edit]

The primary mechanism of soil creation is the weathering of rock. All rock types (igneous rock, metamorphic rock and sedimentary rock) may be broken down into small particles to create soil. Weathering mechanisms are physical weathering, chemical weathering, and biological weathering $[^1][^2][^3]$ Human activities such as excavation, blasting, and waste disposal, may also create soil. Over geologic time, deeply buried soils may be altered by pressure and temperature to become metamorphic or sedimentary rock, and if melted and solidified again, they would complete the geologic cycle by becoming igneous rock. $[^3]$

Physical weathering includes temperature effects, freeze and thaw of water in cracks, rain, wind, impact and other mechanisms. Chemical weathering includes dissolution of matter composing a rock and precipitation in the form of another mineral. Clay minerals, for example can be formed by weathering of feldspar, which is the most common mineral present in igneous rock.

The most common mineral constituent of silt and sand is quartz, also called silica, which has the chemical name silicon dioxide. The reason that feldspar is most common in rocks but silica is more prevalent in soils is that feldspar is much more soluble than silica.

Silt, Sand, and Gravel are basically little pieces of broken rocks.

According to the Unified Soil Classification System, silt particle sizes are in the range of 0.002 mm to 0.075 mm and sand particles have sizes in the range of 0.075 mm to 4.75 mm.

Gravel particles are broken pieces of rock in the size range 4.75 mm to 100 mm. Particles larger than gravel are called cobbles and boulders.^[1]^[2]

Transport



Example soil horizons. a) top soil and colluvium b) mature residual soil c) young residual soil d) weathered rock

Soil deposits are affected by the mechanism of transport and deposition to their location. Soils that are not transported are called residual soils—they exist at the same location as the rock from which they were generated. Decomposed granite is a common example of a residual soil. The common mechanisms of transport are the actions of gravity, ice, water, and wind. Wind blown soils include dune sands and loess. Water carries particles of different size depending on the speed of the water, thus soils transported by water are graded according to their size. Silt and clay may settle out in a lake, and gravel and sand collect at the bottom of a river bed. Wind blown soil deposits (aeolian soils) also tend to be sorted according to their grain size. Erosion at the base of glaciers is powerful enough to pick up large rocks and boulders as well as soil; soils dropped by melting ice can be a well graded mixture of widely varying particle sizes. Gravity on its own may also carry particles down from the top of a mountain to make a pile of soil and boulders at the base; soil deposits transported by gravity are called colluvium.[¹]²]

The mechanism of transport also has a major effect on the particle shape. For example, low velocity grinding in a river bed will produce rounded particles. Freshly fractured colluvium particles often have a very angular shape.

Soil composition

[edit]

Soil mineralogy

Silts, sands and gravels are classified by their size, and hence they may consist of a variety of minerals. Owing to the stability of quartz compared to other rock minerals, quartz is the most common constituent of sand and silt. Mica, and feldspar are other common minerals present in sands and silts.^[1] The mineral constituents of gravel may be more similar to that of the parent rock.

The common clay minerals are montmorillonite or smectite, illite, and kaolinite or kaolin. These minerals tend to form in sheet or plate like structures, with length typically ranging between 10 $?^7$ m and 4x10^{?6} m and thickness typically ranging between 10^{?9} m and 2x10^{?6} m, and they have a relatively large specific surface area. The specific surface area (SSA) is defined as the ratio of the surface area of particles to the mass of the particles. Clay minerals typically have specific surface areas in the range of 10 to 1,000 square meters per gram of solid.^[3] Due to the large surface area available for chemical, electrostatic, and van der Waals interaction, the mechanical behavior of clay minerals is very sensitive to the amount of pore fluid available and the type and amount of dissolved ions in the pore fluid.^{[1}]

The minerals of soils are predominantly formed by atoms of oxygen, silicon, hydrogen, and aluminum, organized in various crystalline forms. These elements along with calcium, sodium, potassium, magnesium, and carbon constitute over 99 per cent of the solid mass of soils.^[1]

Grain size distribution

[edit] Main article: Soil gradation

Soils consist of a mixture of particles of different size, shape and mineralogy. Because the size of the particles obviously has a significant effect on the soil behavior, the grain size and grain size distribution are used to classify soils. The grain size distribution describes the relative proportions of particles of various sizes. The grain size is often visualized in a cumulative distribution graph which, for example, plots the percentage of particles finer than a given size as a function of size. The median grain size, display the specially the hydraulic conductivity, tends to be dominated by the smaller particles, hence, the term "effective size", denoted by displaystyle, p. 10, own defined as the size for which 10% of the particle mass consists of finer particles.

Sands and gravels that possess a wide range of particle sizes with a smooth distribution of particle sizes are called *well graded* soils. If the soil particles in a sample are predominantly in a relatively narrow range of sizes, the sample is *uniformly graded*. If a soil sample has distinct gaps in the gradation curve, e.g., a mixture of gravel and fine sand, with no coarse sand, the sample may be *gap graded*. *Uniformly graded* and *gap graded* soils are both considered to be *poorly graded*. There are many methods for measuring particle-size distribution. The two traditional methods are sieve analysis and hydrometer analysis.

Sieve analysis

[edit]



The size distribution of gravel and sand particles are typically measured using sieve analysis. The formal procedure is described in ASTM D6913-04(2009).[⁸] A stack of sieves with accurately dimensioned holes between a mesh of wires is used to separate the particles into size bins. A known volume of dried soil, with clods broken down to individual particles, is put into the top of a stack of sieves arranged from coarse to fine. The stack of sieves is shaken for a standard period of time so that the particles are sorted into size bins. This method works reasonably well for particles in the sand and gravel size range. Fine particles tend to stick to each other, and hence the sieving process is not an effective method. If there are a lot of fines (silt and clay) present in the soil it may be necessary to run water through the sieves to wash the coarse particles and clods through.

A variety of sieve sizes are available. The boundary between sand and silt is arbitrary. According to the Unified Soil Classification System, a #4 sieve (4 openings per inch) having 4.75 mm opening size separates sand from gravel and a #200 sieve with an 0.075 mm opening separates sand from silt and clay. According to the British standard, 0.063 mm is the boundary between sand and silt, and 2 mm is the boundary between sand and gravel.³]

Hydrometer analysis

[edit]

The classification of fine-grained soils, i.e., soils that are finer than sand, is determined primarily by their Atterberg limits, not by their grain size. If it is important to determine the grain size distribution of fine-grained soils, the hydrometer test may be performed. In the hydrometer tests, the soil particles are mixed with water and shaken to produce a dilute suspension in a glass cylinder, and then the cylinder is left to sit. A hydrometer is used to measure the density of the suspension as a function of time. Clay particles may take several hours to settle past the

depth of measurement of the hydrometer. Sand particles may take less than a second. Stokes' law provides the theoretical basis to calculate the relationship between sedimentation velocity and particle size. ASTM provides the detailed procedures for performing the Hydrometer test.

Clay particles can be sufficiently small that they never settle because they are kept in suspension by Brownian motion, in which case they may be classified as colloids.

Mass-volume relations

[edit]



A phase diagram of soil indicating the masses and volumes of air, solid, water, and voids

There are a variety of parameters used to describe the relative proportions of air, water and solid in a soil. This section defines these parameters and some of their interrelationships.^{[2}]^{[6}] The basic notation is as follows:

kisphissenes wiere and solids in a soil mixture;

displais where the weights of air, water and solids in a soil mixture;

displaysory with be masses of air, water and solids in a soil mixture;

disputions and solids) in a soil mixture;

Note that the weights, W, can be obtained by multiplying the mass, M, by the acceleration due to gravity, g; e.g., kisplaystyle, W_s=M_sg

Specific Gravity is the ratio of the density of one material compared to the density of pure \displaystyle \rho w=1g/cm^3 water (Image not found or type) Inknown

\displaystyle G_s=\frac \rho _s\rho _w

Specific gravity of solids, Image not found or type unknown

Note that specific weight, conventionally denoted by the symbol har by be have have multiplying the density (has favore in by the acceleration due to gravity, he is playety leve unknown

Density, bulk density, or wet density, displaying the pames for the density of the mixture, i.e., the total mass of air, water, solids divided by the total volume of air water and solids (the mass of air is assumed to be zero for practical purposes):

```
\displaystyle \rho =\frac M_s+M_wV_s+V_w+V_a=\frac M_tV_t
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Dry density, disister many of splids divided by the total volume of air water and solids:

\displaystyle \rho _d=\frac M_sV_s+V_w+V_a=\frac M_sV_t

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Buoyant density, he density of the mixture minus the density of water is useful if the soil is submerged under water:

\displaystyle \rho '=\rho \ -\rho _w

where holis is the density of water

Water content, which the soil, drying it out in an oven and re-weighing. Standard procedures are described by ASTM.

\displaystyle w=\frac M_wM_s=\frac W_wW_s

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Void ratio, vois to the volume of voids to the volume of solids:

\displaystyle e=\frac V_vV_s=\frac V_vV_t-V_v=\frac n1-n

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Porosity, this the void ratio:

\displaystyle n=\frac V_vV_t=\frac V_vV_s+V_v=\frac e1+e

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Degree of saturation, displate Sof the volume of water to the volume of voids:

\displaystyle S=\frac V_wV_v

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From the above definitions, some useful relationships can be derived by use of basic algebra.

\displaystyle \rho =\frac (G_s+Se)\rho _w1+e

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\displaystyle \rho =\frac (1+w)G_s\rho _w1+e

Image not found or type unknown

\displaystyle w=\frac SeG_s

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Soil classification

[edit]

Geotechnical engineers classify the soil particle types by performing tests on disturbed (dried, passed through sieves, and remolded) samples of the soil. This provides information about the characteristics of the soil grains themselves. Classification of the types of grains present in a soil does not *clarification needed* account for important effects of the *structure* or *fabric* of the soil, terms that describe compactness of the particles and patterns in the arrangement of particles in a load carrying framework as well as the pore size and pore fluid distributions. Engineering geologists also classify soils based on their genesis and depositional history.

Classification of soil grains

[edit]

In the US and other countries, the Unified Soil Classification System (USCS) is often used for soil classification. Other classification systems include the British Standard BS 5930 and the AASHTO soil classification system.^[3]

Classification of sands and gravels

[edit]

In the USCS, gravels (given the symbol *G*) and sands (given the symbol *S*) are classified according to their grain size distribution. For the USCS, gravels may be given the classification symbol *GW* (well-graded gravel), *GP* (poorly graded gravel), *GM* (gravel with a large amount of silt), or *GC* (gravel with a large amount of clay). Likewise sands may be classified as being *SW*, *SP*, *SM* or *SC*. Sands and gravels with a small but non-negligible amount of fines (5–12%) may be given a dual classification such as *SW-SC*.

Atterberg limits

[edit]

Clays and Silts, often called 'fine-grained soils', are classified according to their Atterberg limits; the most commonly used Atterberg limits are the *liquid limit* (denoted by *LL* or *displaystyle style limit* (denoted by *PL* or *displaystyle limit* (denoted by *SL*).

The liquid limit is the water content at which the soil behavior transitions from a plastic solid to a liquid. The plastic limit is the water content at which the soil behavior transitions from that of a plastic solid to a brittle solid. The Shrinkage Limit corresponds to a water content below which the soil will not shrink as it dries. The consistency of fine grained soil varies in proportional to the water content in a soil.

As the transitions from one state to another are gradual, the tests have adopted arbitrary definitions to determine the boundaries of the states. The liquid limit is determined by measuring the water content for which a groove closes after 25 blows in a standard test.[⁹][[]*clarificatio* Alternatively, a fall cone test apparatus may be used to measure the liquid limit. The undrained shear strength of remolded soil at the liquid limit is approximately 2 kPa.[⁴][¹⁰] The plastic limit is the water content below which it is not possible to roll by hand the soil into 3 mm diameter cylinders. The soil cracks or breaks up as it is rolled down to this diameter. Remolded soil at the plastic limit is quite stiff, having an undrained shear strength of the order of about 200 kPa.[⁴][¹⁰]

The *plasticity index* of a particular soil specimen is defined as the difference between the liquid limit and the plastic limit of the specimen; it is an indicator of how much water the soil particles in the specimen can absorb, and correlates with many engineering properties like permeability, compressibility, shear strength and others. Generally, the clay having high plasticity have lower permeability and also they are also difficult to be compacted.

Classification of silts and clays

[edit]

According to the Unified Soil Classification System (USCS), silts and clays are classified by plotting the values of their plasticity index and liquid limit on a plasticity chart. The A-Line on the chart separates clays (given the USCS symbol *C*) from silts (given the symbol *M*). LL=50% separates high plasticity soils (given the modifier symbol *H*) from low plasticity soils (given the modifier symbol *L*). A soil that plots above the A-line and has LL>50% would, for example, be classified as *CH*. Other possible classifications of silts and clays are *ML*, *CL* and *MH*. If the Atterberg limits plot in the"hatched" region on the graph near the origin, the soils are given the dual classification 'CL-ML'.

Indices related to soil strength

[edit]

Liquidity index

[edit]

The effects of the water content on the strength of saturated remolded soils can be quantified by the use of the *liquidity index*, *LI*:

```
\displaystyle LI=\frac w-PLLL-PL
```

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When the LI is 1, remolded soil is at the liquid limit and it has an undrained shear strength of about 2 kPa. When the soil is at the plastic limit, the LI is 0 and the undrained shear strength is about 200 kPa.[⁴][¹¹]

Relative density

[edit]

The density of sands (cohesionless soils) is often characterized by the relative density, displaystyle,

\displaystyle D_r=\frac e_max-ee_max-e_min100\%

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where: \displayshe maximum void ratio" corresponding to a very loose state, \displayshe minum "minimum void ratio" corresponding to a very dense state and \displayshe woid ratio. Methods used to calculate relative density are defined in ASTM D4254-00(2006).[¹²]

Thus if \displaystyle. The stand or gravel is very dense, and if \displaystyle. Stand or gravel is very dense, and if \displaystyle.

Seepage: steady state flow of water

[edit] This section is an excerpt from Seepage.[edit]



A cross section showing the water table varying with surface topography as well as a perched water table

In soil mechanics, seepage is the movement of water through soil. If fluid pressures in a soil deposit are uniformly increasing with depth according to displaystyle. Where displaystyle where table, then hydrostatic conditions will prevail and the fluids will *not* be flowing through the soil. However, if the water table is sloping or there is a perched water table as indicated in the accompanying sketch, then seepage will occur. For steady state seepage, the seepage velocities are not varying with time. If the water tables are changing levels with time, or if the soil is in the process of consolidation, then steady state conditions do not apply.

Effective stress and capillarity: hydrostatic conditions

[edit]



Spheres immersed in water, reducing effective stress

Main article: Effective stress

To understand the mechanics of soils it is necessary to understand how normal stresses and shear stresses are shared by the different phases. Neither gas nor liquid provide significant resistance to shear stress. The shear resistance of soil is provided by friction and interlocking of the particles. The friction depends on the intergranular contact stresses between solid particles. The normal stresses, on the other hand, are shared by the fluid and the particles.[⁷] Although the pore air is relatively compressible, and hence takes little normal stress in most geotechnical problems, liquid water is relatively incompressible and if the voids are saturated with water, the pore water must be squeezed out in order to pack the particles closer together.

The principle of effective stress, introduced by Karl Terzaghi, states that the effective stress ?' (i.e., the average intergranular stress between solid particles) may be calculated by a simple subtraction of the pore pressure from the total stress:

```
\displaystylersigma, =\sigma -u\,
```

where ? is the total stress and *u* is the pore pressure. It is not practical to measure ?' directly, so in practice the vertical effective stress is calculated from the pore pressure and vertical total stress. The distinction between the terms pressure and stress is also important. By definition, pressure at a point is equal in all directions but stresses at a point can be different in different directions. In soil mechanics, compressive stresses and pressures are considered to be positive and tensile stresses are considered to be negative, which is different from the solid mechanics sign convention for stress.

Total stress

[edit]

For level ground conditions, the total vertical stress at a point, displayere any here weight of everything above that point per unit area. The vertical stress beneath a uniform surface layer with density diaped thicker as displayered and lenown

```
\displaystyle,\sigma,_v=\rho gH=\gamma H
```

where *displayageleration* due to gravity, and *displayet pieve interview of the overlying layer*. If there are multiple layers of soil or water above the point of interest, the vertical stress may be calculated by summing the product of the unit weight and thickness of all of the overlying layers. Total stress increases with increasing depth in proportion to the density of the overlying soil.

It is not possible to calculate the horizontal total stress in this way. Lateral earth pressures are addressed elsewhere.

Pore water pressure

[edit] Main article: Pore water pressure

Hydrostatic conditions

[edit]



Water is drawn into a small tube by surface tension. Water pressure, u, is negative above and positive below the free water surface.

If the soil pores are filled with water that is not flowing but is static, the pore water pressures will be hydrostatic. The water table is located at the depth where the water pressure is equal to the atmospheric pressure. For hydrostatic conditions, the water pressure increases linearly with depth below the water table:

\displaystyleyu=\rhovn_wgz_w

where hoising the water, and hoising the part of the water table.

Capillary action

[edit]

Due to surface tension, water will rise up in a small capillary tube above a free surface of water. Likewise, water will rise up above the water table into the small pore spaces around the soil particles. In fact the soil may be completely saturated for some distance above the water table. Above the height of capillary saturation, the soil may be wet but the water content will decrease with elevation. If the water in the capillary zone is not moving, the water pressure obeys the equation of hydrostatic equilibrium, <code>displaystyle.butnote.totabove</code> the water table. Hence, hydrostatic water pressures are negative above the water table. The thickness of the zone of capillary saturation depends on the pore size, but typically, the heights vary between a centimeter or so for coarse sand to tens of meters for a silt or clay.^[3] In fact the pore space of soil is a uniform fractal e.g. a set of uniformly distributed D-dimensional fractals of average linear size L. For the clay soil it has been found that L=0.15 mm and D=2.7. [¹³]

The surface tension of water explains why the water does not drain out of a wet sand castle or a moist ball of clay. Negative water pressures make the water stick to the particles and pull the

particles to each other, friction at the particle contacts make a sand castle stable. But as soon as a wet sand castle is submerged below a free water surface, the negative pressures are lost and the castle collapses. Considering the effective stress equation, displaystyle sigma 'a sigma -u, pressure is negative, the effective stress may be positive, even on a free surface (a surface where the total normal stress is zero). The negative pore pressure pulls the particles together and causes compressive particle to particle contact forces. Negative pore pressures in clayey soil can be much more powerful than those in sand. Negative pore pressures explain why clay soils shrink when they dry and swell as they are wetted. The swelling and shrinkage can cause major distress, especially to light structures and roads.[¹⁴]

Later sections of this article address the pore water pressures for seepage and consolidation problems.

Water at particle contacts

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• Intergranular contact force due to surface tension

Image not found or type unknown Intergranular contact force due to surface tension Shrinkage caused by drying

0

Image not found or type unknown Shrinkage caused by drying

Consolidation: transient flow of water

[edit] Main article: Consolidation (soil)



Consolidation analogy. The piston is supported by water underneath and a spring. When a load is applied to the piston, water pressure increases to support the load. As the water slowly leaks through the small hole, the load is transferred from the water pressure to the spring force.

Consolidation is a process by which soils decrease in volume. It occurs when stress is applied to a soil that causes the soil particles to pack together more tightly, therefore reducing volume. When this occurs in a soil that is saturated with water, water will be squeezed out of the soil. The time required to squeeze the water out of a thick deposit of clayey soil layer might be years. For a layer of sand, the water may be squeezed out in a matter of seconds. A building foundation or construction of a new embankment will cause the soil below to consolidate and this will cause settlement which in turn may cause distress to the building or embankment. Karl Terzaghi developed the theory of one-dimensional consolidation which enables prediction of the amount of settlement and the time required for the settlement to occur.^{[15}] Afterwards, Maurice Biot fully developed the three-dimensional soil consolidation theory, extending the one-dimensional model previously developed by Terzaghi to more general hypotheses and introducing the set of basic equations of Poroelasticity.^{[7}] Soils are tested with an oedometer test to determine their compression index and coefficient of consolidation.

When stress is removed from a consolidated soil, the soil will rebound, drawing water back into the pores and regaining some of the volume it had lost in the consolidation process. If the stress is reapplied, the soil will re-consolidate again along a recompression curve, defined by the recompression index. Soil that has been consolidated to a large pressure and has been subsequently unloaded is considered to be *overconsolidated*. The maximum past vertical effective stress is termed the *preconsolidation stress*. A soil which is currently experiencing the maximum past vertical effective stress is said to be *normally consolidated*. The *overconsolidation ratio*, (OCR) is the ratio of the maximum past vertical effective stress to the current vertical effective stress. The OCR is significant for two reasons: firstly, because the compressibility of normally consolidated soil is significantly larger than that for overconsolidated soil, and secondly, the shear behavior and dilatancy of clayey soil are related to the OCR through critical state soil mechanics; highly overconsolidated clayey soils are dilatant, while normally consolidated soils tend to be contractive.[²][³][⁴]

Shear behavior: stiffness and strength

[edit] Main article: shear strength (soil)



Typical stress strain curve for a drained dilatant soil

The shear strength and stiffness of soil determines whether or not soil will be stable or how much it will deform. Knowledge of the strength is necessary to determine if a slope will be stable, if a building or bridge might settle too far into the ground, and the limiting pressures on a retaining wall. It is important to distinguish between failure of a soil element and the failure of a geotechnical structure (e.g., a building foundation, slope or retaining wall); some soil elements may reach their peak strength prior to failure of the structure. Different criteria can be used to define the "shear strength" and the "yield point" for a soil element from a stress–strain curve. One may define the peak shear strength as the peak of a stress–strain curve, or the shear strength at critical state as the value after large strains when the shear resistance levels off. If the stress–strain curve does not stabilize before the end of shear strength test, the "strength" is sometimes considered to be the shear resistance at 15–20% strain.[¹⁴] The shear strength of soil depends on many factors including the effective stress and the void ratio.

The shear stiffness is important, for example, for evaluation of the magnitude of deformations of foundations and slopes prior to failure and because it is related to the shear wave velocity. The slope of the initial, nearly linear, portion of a plot of shear stress as a function of shear strain is called the shear modulus

Friction, interlocking and dilation



Angle of repose

Soil is an assemblage of particles that have little to no cementation while rock (such as sandstone) may consist of an assembly of particles that are strongly cemented together by chemical bonds. The shear strength of soil is primarily due to interparticle friction and therefore, the shear resistance on a plane is approximately proportional to the effective normal stress on that plane.^[3] The angle of internal friction is thus closely related to the maximum stable slope angle, often called the angle of repose.

But in addition to friction, soil derives significant shear resistance from interlocking of grains. If the grains are densely packed, the grains tend to spread apart from each other as they are subject to shear strain. The expansion of the particle matrix due to shearing was called dilatancy by Osborne Reynolds.^[11] If one considers the energy required to shear an assembly of particles there is energy input by the shear force, T, moving a distance, x and there is also energy input by the normal force, N, as the sample expands a distance, y.^[11] Due to the extra energy required for the particles to dilate against the confining pressures, dilatant soils have a greater peak strength than contractive soils. Furthermore, as dilative soil grains dilate, they become looser (their void ratio increases), and their rate of dilation decreases until they reach a critical void ratio. Contractive soils become denser as they shear, and their rate of contraction decreases until they reach a critical void ratio.



A critical state line separates the dilatant and contractive states for soil.

The tendency for a soil to dilate or contract depends primarily on the confining pressure and the void ratio of the soil. The rate of dilation is high if the confining pressure is small and the void ratio is small. The rate of contraction is high if the confining pressure is large and the void ratio is large. As a first approximation, the regions of contraction and dilation are separated by the critical state line.

Failure criteria

After a soil reaches the critical state, it is no longer contracting or dilating and the shear stress on the failure plane displayer mined by the effective normal stress on the failure plane displayer and critical state friction angle displayer by the crit'

```
\displaystyle.\tau crit=\sigma _n'\tan \phi _crit'\
```

The peak strength of the soil may be greater, however, due to the interlocking (dilatancy) contribution. This may be stated:

\displaystyle \tau peak=\sigma _n'\tan \phi _peak'\

where $\underset{\text{mage not found or type in every, the origination of type in every, the origination of type in every, the origin requires care. The peak strength will not be mobilized everywhere at the same time in a practical problem such as a foundation, slope or retaining wall. The critical state friction angle is not nearly as variable as the peak friction angle and hence it can be relied upon with confidence.[³][⁴][¹¹]$

Not recognizing the significance of dilatancy, Coulomb proposed that the shear strength of soil may be expressed as a combination of adhesion and friction components:[¹¹]

\displaystyle \tau _f=c'+\sigma _f'\tan \phi '\, Image not found or type unknown

It is now known that the ^{\displaydi}

Structure, fabric, and chemistry

[edit]

In addition to the friction and interlocking (dilatancy) components of strength, the structure and fabric also play a significant role in the soil behavior. The structure and fabric include factors such as the spacing and arrangement of the solid particles or the amount and spatial distribution of pore water; in some cases cementitious material accumulates at particle-particle contacts. Mechanical behavior of soil is affected by the density of the particles and their structure or arrangement of the particles as well as the amount and spatial distribution of fluids present (e.g., water and air voids). Other factors include the electrical charge of the particles, chemistry of pore water, chemical bonds (i.e. cementation -particles connected through a solid substance such as recrystallized calcium carbonate) [¹][¹⁶]

Drained and undrained shear

[edit]



Moist sand along the shoreline is originally densely packed by the draining water. Foot pressure on the sand causes it to dilate *(see: Reynolds dilatancy)*, drawing water from the surface into the pores.

The presence of nearly incompressible fluids such as water in the pore spaces affects the ability for the pores to dilate or contract.

If the pores are saturated with water, water must be sucked into the dilating pore spaces to fill the expanding pores (this phenomenon is visible at the beach when apparently dry spots form around feet that press into the wet sand). [clarification needed]

Similarly, for contractive soil, water must be squeezed out of the pore spaces to allow contraction to take place.

Dilation of the voids causes negative water pressures that draw fluid into the pores, and contraction of the voids causes positive pore pressures to push the water out of the pores. If the rate of shearing is very large compared to the rate that water can be sucked into or squeezed out of the dilating or contracting pore spaces, then the shearing is called *undrained shear*, if the shearing is slow enough that the water pressures are negligible, the shearing is called *drained shear*. During undrained shear, the water pressure u changes depending on volume change tendencies. From the effective stress equation, the change in u directly effects the effective stress by the equation:

\displaystyle_\sigma,'=\sigma -u\,

and the strength is very sensitive to the effective stress. It follows then that the undrained shear strength of a soil may be smaller or larger than the drained shear strength depending upon whether the soil is contractive or dilative.

Shear tests

[edit]

Strength parameters can be measured in the laboratory using direct shear test, triaxial shear test, simple shear test, fall cone test and (hand) shear vane test; there are numerous other devices and variations on these devices used in practice today. Tests conducted to characterize the strength and stiffness of the soils in the ground include the Cone penetration test and the Standard penetration test.

Other factors

[edit]

The stress–strain relationship of soils, and therefore the shearing strength, is affected by:[¹⁷]

- 1. *soil composition* (basic soil material): mineralogy, grain size and grain size distribution, shape of particles, pore fluid type and content, ions on grain and in pore fluid.
- 2. *state* (initial): Defined by the initial void ratio, effective normal stress and shear stress (stress history). State can be described by terms such as: loose, dense, overconsolidated, normally consolidated, stiff, soft, contractive, dilative, etc.
- 3. *structure*: Refers to the arrangement of particles within the soil mass; the manner in which the particles are packed or distributed. Features such as layers, joints, fissures, slickensides, voids, pockets, cementation, etc., are part of the structure. Structure of soils is described by terms such as: undisturbed, disturbed, remolded, compacted, cemented; flocculent, honey-combed, single-grained; flocculated, deflocculated; stratified, layered, laminated; isotropic and anisotropic.
- 4. Loading conditions: Effective stress path drained, undrained, and type of loading magnitude, rate (static, dynamic), and time history (monotonic, cyclic).

Applications

[edit]

Lateral earth pressure

Main article: Lateral earth pressure

Lateral earth stress theory is used to estimate the amount of stress soil can exert perpendicular to gravity. This is the stress exerted on retaining walls. A lateral earth stress coefficient, K, is defined as the ratio of lateral (horizontal) effective stress to vertical effective stress for cohesionless soils (K=?'_h/?'_v). There are three coefficients: at-rest, active, and passive. At-rest stress is the lateral stress in the ground before any disturbance takes place. The active stress state is reached when a wall moves away from the soil under the influence of lateral stress, and results from shear failure due to reduction of lateral stress. The passive stress state is reached when a wall is pushed into the soil far enough to cause shear failure within the mass due to increase of lateral stress. There are many theories for estimating lateral earth stress; some are empirically based, and some are analytically derived.

Bearing capacity

[edit] Main article: Bearing capacity

The bearing capacity of soil is the average contact stress between a foundation and the soil which will cause shear failure in the soil. Allowable bearing stress is the bearing capacity divided by a factor of safety. Sometimes, on soft soil sites, large settlements may occur under loaded foundations without actual shear failure occurring; in such cases, the allowable bearing stress is determined with regard to the maximum allowable settlement. It is important during construction and design stage of a project to evaluate the subgrade strength. The California Bearing Ratio (CBR) test is commonly used to determine the suitability of a soil as a subgrade for design and construction. The field Plate Load Test is commonly used to predict the deformations and failure characteristics of the soil/subgrade and modulus of subgrade reaction (ks). The Modulus of subgrade reaction (ks) is used in foundation design, soil-structure interaction studies and design of highway pavements.[[]*citation needed*]

Slope stability



Simple slope slip section

Main article: Slope stability

The field of slope stability encompasses the analysis of static and dynamic stability of slopes of earth and rock-fill dams, slopes of other types of embankments, excavated slopes, and natural slopes in soil and soft rock.^[18]

As seen to the right, earthen slopes can develop a cut-spherical weakness zone. The probability of this happening can be calculated in advance using a simple 2-D circular analysis package.[¹⁹] A primary difficulty with analysis is locating the most-probable slip plane for any given situation.[²⁰] Many landslides have been analyzed only after the fact. Landslides vs. Rock strength are two factors for consideration.

Recent developments

[edit]

A recent finding in soil mechanics is that soil deformation can be described as the behavior of a dynamical system. This approach to soil mechanics is referred to as Dynamical Systems based Soil Mechanics (DSSM). DSSM holds simply that soil deformation is a Poisson process in which particles move to their final position at random shear strains.

The basis of DSSM is that soils (including sands) can be sheared till they reach a steady-state condition at which, under conditions of constant strain-rate, there is no change in shear stress, effective confining stress, and void ratio. The steady-state was formally defined[²¹] by Steve J. Poulos Archived 2020-10-17 at the Wayback Machine an associate professor at the Soil Mechanics Department of Harvard University, who built off a hypothesis that Arthur Casagrande was formulating towards the end of his career. The steady state condition is not the same as the "critical state" condition. It differs from the critical state in that it specifies a statistically constant structure at the steady state. The steady-state values are also very slightly dependent on the strain-rate.

Many systems in nature reach steady states, and dynamical systems theory describes such systems. Soil shear can also be described as a dynamical system.^{[22}]^{[23}] The physical basis of the soil shear dynamical system is a Poisson process in which particles move to the steady-state at random shear strains.^{[24}] Joseph^{[25}] generalized this—particles move to their final

position (not just steady-state) at random shear-strains. Because of its origins in the steady state concept, DSSM is sometimes informally called "Harvard soil mechanics."

DSSM provides for very close fits to stress-strain curves, including for sands. Because it tracks conditions on the failure plane, it also provides close fits for the post failure region of sensitive clays and silts something that other theories are not able to do. Additionally DSSM explains key relationships in soil mechanics that to date have simply been taken for granted, for example, why normalized undrained peak shear strengths vary with the log of the overconsolidation ratio and why stress-strain curves normalize with the initial effective confining stress; and why in one-dimensional consolidation the void ratio must vary with the log of the effective vertical stress, why the end-of-primary curve is unique for static load increments, and why the ratio of the creep value C? to the compression index Cc must be approximately constant for a wide range of soils.²⁶]

See also

[edit]

- Critical state soil mechanics
- Earthquake engineering
- Engineering geology
- Geotechnical centrifuge modeling
- Geotechnical engineering
- Geotechnical engineering (Offshore)
- Geotechnics
- · Hydrogeology, aquifer characteristics closely related to soil characteristics
- International Society for Soil Mechanics and Geotechnical Engineering
- Rock mechanics
- Slope stability analysis

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External links

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- Media related to Soil mechanics at Wikimedia Commons
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- **e**

Soil science

- History
- Index
- Pedology
- Edaphology
- Soil biology
- Soil microbiology
- Soil zoology

Main fields

- Soil ecologySoil physics
- Soil mechanics
- Soil chemistry
- Environmental soil science
- Agricultural soil science



- Soil
- Pedosphere
 - Soil morphology
 - Pedodiversity
 - Soil formation
- \circ Soil erosion
- Soil contamination
- Soil retrogression and degradation
- Soil compaction
 - Soil compaction (agriculture)
- Soil sealing
- Soil salinity
 - Alkali soil
- Soil pH
 - Soil acidification
- Soil health
- Soil life

Soil topics

- Soil biodiversity
- Soil quality
- Soil value
- Soil fertility
- Soil resilience
- \circ Soil color
- Soil texture
- Soil structure
 - $\circ\,$ Pore space in soil
 - Pore water pressure
- Soil crust
- Soil horizon
- Soil biomantle
- Soil carbon
- Soil gas
 - Soil respiration
- Soil organic matter
- Soil moisture
 - Soil water (retention)

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Soil classification

- Acrisols
- Alisols
- Andosols
- Anthrosols
- Arenosols
- Calcisols
- Cambisols
- Chernozem
- Cryosols
- Durisols
- Ferralsols
- Fluvisols
- Gleysols • Gypsisols
- World Reference Base for Soil
- Resources (1998–)
- Kastanozems • Leptosols

• Histosol

- Lixisols
- Luvisols
- Nitisols
- Phaeozems
- Planosols
- Plinthosols
- Podzols
- Regosols
- Retisols
- Solonchaks
- Solonetz
- Stagnosol
- Technosols
- Umbrisols
- Vertisols
- Alfisols
- Andisols
- Aridisols
- Entisols
- Gelisols
- **USDA** soil
- Incontingle
- Histosols

- Soil conservation
- Soil management
- Soil guideline value
- Soil survey
- Soil test

Applications

- Soil governance
 Soil value
- Soil value
 Soil salinity control
- Soli sainity control
 Erosion control
- Agroecology
- Liming (soil)
- Geology
- Geochemistry
- Petrology
- Geomorphology
- Geotechnical engineering

Related • Hydrology

fields

- HydrogeologyBiogeography
- Earth materials
- Archaeology
- Agricultural science
 - Agrology
- Australian Society of Soil Science Incorporated
- Canadian Society of Soil Science
- Central Soil Salinity Research Institute (India)
- German Soil Science Society
- Indian Institute of Soil Science
- $\circ\,$ International Union of Soil Sciences

Societies, Initiatives

- International Year of Soil
 National Society of Consulting Soil Scientists (US)
- OPAL Soil Centre (UK)
- Soil Science Society of Poland
- Soil and Water Conservation Society (US)
- Soil Science Society of America
- World Congress of Soil Science

- Acta Agriculturae Scandinavica B
- Journal of Soil and Water Conservation

Scientific journals

- Plant and Soil
 - Pochvovedenie
 - Soil Research
 - Soil Science Society of America Journal
 - Land use
 - \circ Land conversion
 - Land management
 - Vegetation
- See also
- Infiltration (hydrology)
 - Groundwater
 - Crust (geology)
 - Impervious surface/Surface runoff
 - Petrichor
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- Category soil science
- Eist of soil scientists
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Geotechnical engineering

Offshore geotechnical engineering

	• Cone penetration test
	• Geo-electrical sounding
	• Permeability test
	 Load test Static Dynamic Statnamic
	 Pore pressure measurement Piezometer Well
	• Ram sounding
	• Control drilling
	• Kotary-pressure sounding
	• Kotary weight sounding
	• Sample series
Field (<i>in situ</i>)	• Screw plate test
	 Deformation monitoring

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- Incline und or type unknown
 Incline or type unknown
 Settlement recordings
- Shear vane test
- Standard penetration test
- Total sounding
- $\circ \overset{\text{Image pot found or type unknown}}{\square \text{Trial pit}}$
- Visible bedrock
- Nuclear densometer test
- Exploration geophysics
- Crosshole sonic logging

Investigation and instrumentation

Types	 Clay Silt Sand Gravel Peat Loam Loess
Properties	 Hydraulic conductivity Water content Void ratio Bulk density Thixotropy Reynolds' dilatancy Angle of repose Friction angle Cohesion Porosity Permeability Specific storage Shear strength Sensitivity

Soil

\circ	Topography	. /
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- Vegetation
- Terrain

• Topsoil

Natural features • Vopson • Water table

- Bedrock
- Subgrade
- Subsoil
- Shoring structures
 - Retaining walls
 - Gabion
 - Ground freezing
 - Mechanically stabilized earth
 - Pressure grouting
 - Slurry wall
 - Soil nailing
 - Tieback
- Land development
- Landfill
- Excavation
- Trench
- Embankment
- Cut
- Causeway
- Terracing
- Cut-and-cover
- $\circ~$ Cut and fill
- Fill dirt
- Grading
- Land reclamation
- Track bed
- Erosion control
- Earth structure
- Expanded clay aggregate
- Crushed stone
- Geosynthetics
 - Geotextile
 - Geomembrane
 - $\circ\,$ Geosynthetic clay liner
 - Cellular confinement
- Infiltration

Foundations

Shallow
 Deep

Structures (Interaction)

Earthworks

	Forces	 Effective stress Pore water pressure Lateral earth pressure Overburden pressure Preconsolidation pressure
Mechanics	Phenomena/ problems	 Permafrost Frost heaving Consolidation Compaction Compaction Earthquake Response spectrum Seismic hazard Shear wave Landslide analysis Stability analysis Mitigation Classification Sliding criterion Slab stabilisation Bearing capacity * Stress distribution in soil

	 SEEP2D
	 STABL
Numerical analysis	○ SVFlux
software	 SVSlope
	 UTEXAS
	 Plaxis

- Geology
- Geochemistry
- Petrology
- Earthquake engineering
- Geomorphology
- Soil science

Related fields

- Hydrology
- Hydrogeology
- Biogeography
- Earth materials
- Archaeology
- Agricultural science
 - Agrology
- Germany
- United States
- France

Authority control databases: National Batt of Batt Wikidata

- Japan
- Czech Republic
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About foundation

mage rot found or type unknown

Look up *foundation* or *foundations* in Wiktionary, the free dictionary.

Foundation(s) or The Foundation(s) may refer to:

Common uses

- Foundation (cosmetics), a skin-coloured makeup cream applied to the face
- Foundation (engineering), the element of a structure which connects it to the ground, and transfers loads from the structure to the ground
- Foundation (evidence), a legal term

- Foundation (nonprofit), a type of charitable organization
 - Foundation (United States law), a type of charitable organization in the U.S.
 - Private foundation, a charitable organization that might not qualify as a public charity by government standards

Arts, entertainment, and media

[edit]

Film and TV

[edit]

- The Foundation, a film about 1960s-1970s Aboriginal history in Sydney, featuring Gary Foley
- The Foundation (1984 TV series), a Hong Kong series
- The Foundation (Canadian TV series), a 2009–2010 Canadian sitcom
- "The Foundation" (Seinfeld), an episode
- Foundation (TV series), an Apple TV+ series adapted from Isaac Asimov's novels

Games

[edit]

- Foundation (video game), a city-building game (2025)
- Foundation, an Amiga video game
- The Foundation, a character in 2017 game Fortnite Battle Royale

Literature

- Foundation (book series), a series of science fiction books by Isaac Asimov
 - Foundation (Asimov novel), the first book in Asimov's series, published in 1951
- Foundation (b-boy book), by Joseph G. Schloss
- Foundation (Lackey novel), a 2008 fantasy novel by Mercedes Lackey

Music

[edit]

- The Foundations, a British soul group
- Foundations (EP), by Serj Tankian

Albums

[edit]

- Foundation (Brand Nubian album)
- Foundation (Breakage album)
- Foundation (Doc Watson album)
- Foundation (Magnum album)
- Foundation (M.O.P. album)
- Foundation, a 1997 compilation album by Die Krupps
- The Foundation (Geto Boys album)
- The Foundation (Pep Love album), 2005
- The Foundation (Zac Brown Band album)
- The Foundations (album), by 4 Corners

Songs

[edit]

- $\circ\,$ "Foundation", a 1983 song by Spandau Ballet from the album True
- "Foundation", a 1998 song by Brand Nubian from the eponymous album Foundation
- "Foundation", a 2009 song by M.O.P. from the eponymous album Foundation
- "Foundation", a 2010 song by Breakage from the eponymous album Foundation
- "Foundation", a 2015 song by Years & Years from Communion
- "Foundations" (song), by Kate Nash
- "The Foundation" (song), by Xzibit

Other uses in arts, entertainment, and media

- Foundation The International Review of Science Fiction, a literary journal
- The Foundation Trilogy (BBC Radio), a radio adaption of Asimov's series
- The SCP Foundation, a fictional organization that is often referred to in-universe as "The Foundation"

Education

[edit]

- Foundation degree, a British academic qualification
- Foundation school, a type of school in England and Wales
- Foundation Stage, a stage of education for children aged 3 to 5 in England
- University Foundation Programme, a British university entrance course

Science and technology

[edit]

- Foundation (framework), a free collection of tools for creating websites and web applications by ZURB
- Foundation Fieldbus, a communications system
- Foundation Kit, an Apple API

Companies

[edit]

• Foundation Medicine, a genomic profiling company

See also

[edit]

- All pages with titles beginning with *Foundation*
- All pages with titles beginning with The Foundation
- Foundations of mathematics, theory of mathematics

Disambiguation icon

This disambiguation page lists articles associated with the title **Foundation**.

If an internal link led you here, you may wish to change the link to point directly to the intended article.

About Cook County

Photo

Image not found or type unknown **Photo**

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Photo

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Things To Do in Cook County

Photo Image not found or type unknown Sand Ridge Nature Center 4.8 (96) Photo

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River Trail Nature Center	
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Driving Directions in Cook County

Driving Directions From Palmisano (Henry) Park to

Driving Directions From Lake Katherine Nature Center and Botanic Gardens to

Driving Directions From Navy Pier to

https://www.google.com/maps/dir/Navy+Pier/United+Structural+Systems+of+Illinois%2C+I 87.6050944,14z/data=!3m1!4b1!4m14!4m13!1m5!1m1!1sunknown!2m2!1d-87.6050944!2d41.8918633!1m5!1m1!1sChIJ-wSxDtinD4gRiv4kY3RRh9U!2m2!1d-88.1396465!2d42.0637725!3e0

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Reviews for



Jeffery James



Very happy with my experience. They were prompt and followed through, and very helpful in fixing the crack in my foundation.

1						
	lmage nøt	found	or ty	ype	unkno	wn
	\sim					

Sarah McNeily



USS was excellent. They are honest, straightforward, trustworthy, and conscientious. They thoughtfully removed the flowers and flower bulbs to dig where they needed in the yard, replanted said flowers and spread the extra dirt to fill in an area of the yard. We've had other services from different companies and our yard was really a mess after. They kept the job site meticulously clean. The crew was on time and friendly. I'd recommend them any day! Thanks to Jessie and crew.



Jim de Leon (5)

It was a pleasure to work with Rick and his crew. From the beginning, Rick listened to my concerns and what I wished to accomplish. Out of the 6 contractors that quoted the project, Rick seemed the MOST willing to accommodate my wishes. His pricing was definitely more than fair as well. I had 10 push piers installed to stabilize and lift an addition of my house. The project commenced at the date that Rick had disclosed initially and it was completed within the same time period expected (based on Rick's original assessment). The crew was well informed, courteous, and hard working. They were not loud (even while equipment was being utilized) and were well spoken. My neighbors were very impressed on how polite they were when they entered / exited my property (saying hello or good morning each day when they crossed paths). You can tell they care about the customer concerns. They ensured that the property would be put back as clean as possible by placing MANY sheets of plywood down prior to excavating. They compacted the dirt back in the holes extremely well to avoid large stock piles of soils. All the while, the main office was calling me to discuss updates and expectations of completion. They provided waivers of lien, certificates of insurance, properly acquired permits, and JULIE

locates. From a construction background, I can tell you that I did not see any flaws in the way they operated and this an extremely professional company. The pictures attached show the push piers added to the foundation (pictures 1, 2 & 3), the amount of excavation (picture 4), and the restoration after dirt was placed back in the pits and compacted (pictures 5, 6 & 7). Please notice that they also sealed two large cracks and steel plated these cracks from expanding further (which you can see under my sliding glass door). I, as well as my wife, are extremely happy that we chose United Structural Systems for our contractor. I would happily tell any of my friends and family to use this contractor should the opportunity arise!



Chris Abplanalp



USS did an amazing job on my underpinning on my house, they were also very courteous to the proximity of my property line next to my neighbor. They kept things in order with all the dirt/mud they had to excavate. They were done exactly in the timeframe they indicated, and the contract was very details oriented with drawings of what would be done. Only thing that would have been nice, is they left my concrete a little muddy with boot prints but again, all-in-all a great job



Dave Kari (5)

What a fantastic experience! Owner Rick Thomas is a trustworthy professional. Nick and the crew are hard working, knowledgeable and experienced. I interviewed every company in the area, big and small. A homeowner never wants to hear that they have foundation issues. Out of every company, I trusted USS the most, and it paid off in the end. Highly recommend.

Harnessing Infrared Thermography for Hidden Moisture View GBP

Frequently Asked Questions

What is infrared thermography and how does it work?**

** Infrared thermography is non-destructive testing technique uses specialized cameras (thermal imagers) capable detect radiation invisible human eye emitted objects due temperature differences surface areas identify hidden issues like moisture intrusion foundational structures without damaging them during inspection process providing visual

evidence problem areas.

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** Why use IR thermography instead traditional methods?

** Unlike conventional methods which may require drilling holes walls floors potentially causing further damage expensive repairs IR technology allows large surfaces inspected quickly noninvasively pinpoint exact locations problem enabling targeted repairs reducing overall cost time while minimizing disruption homeowner's daily life.

 ** How accurate is IR thermography moisture detection? **

** When performed trained professionals familiar underlying principles equipment used results obtained via IR thermography highly reliable clearly illustrating affected regions allowing accurate diagnosis extent severity moisture issues presence other related problems such mold growth structural weaknesses etc.

**

** Can IR thermography detect moisture behind different materials?

** Yes IR cameras sensitive enough pick up temperature variations caused presence excess moisture various building material surfaces including stucco drywall concrete block wood siding helping locate issues even if visually concealed ensuring thorough assessment entire foundation structure not just visible areas making technology versatile effective tool residential foundation repair services

United Structural Systems of Illinois, Inc

Phone : +18473822882

City : Hoffman Estates

State : IL

Zip : 60169

Address : 2124 Stonington Ave

Google Business Profile

Company Website : https://www.unitedstructuralsystems.com/

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